

Orego

Climate Interpolation Options

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# $NASA/NCEAS/iPlant \ Update$

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#### Climate Aided Interpolation (a brief review)

Theoretical foundation:

- 1. temperature anomalies are strongly correlated out to distances of the order of 1000 km (Hansen and Lebedeff, 1987)
- anomalies are relatively free of the considerable topography-forced spatial variability (Willmott & Robeson, 1995)
- spatial variability within the climatology accounts for most of the temporal between-station variability (Willmott & Robeson, 1995)

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Proposal	Oregon	Precipitation	Climate Interpolation Option

#### Climate Aided Interpolation (a brief review)

CAI

- Di Luzio, M., Johnson, G. L., Daly, C., Eischeid, J. K., & Arnold, J. G. (2008). Constructing Retrospective Gridded Daily Precipitation and Temperature Datasets for the Conterminous United States. *Journal of Applied Meteorology and Climatology*, 47(2), 475497.
- 2. Hunter, R. D., & Meentemeyer, R. K. (2005). Climatologically aided mapping of daily precipitation and temperature. *Journal of Applied Meteorology*, 44(10), 15011510.
- 3. Perry, M., Hollis, D., Perry, M., & Hollis, D. (2005). The generation of monthly gridded datasets for a range of climatic variables over the UK, The generation of monthly gridded datasets for a range of climatic variables over the UK. International Journal of Climatology, *International Journal of Climatology*, 25, 25(8, 8), 1041, 10411054, 1054. doi:10.1002/joc.1161, 10.1002/joc.1161
- Willmott, C. J., & Robeson, S. M. (1995). Climatologically aided interpolation (CAI) of terrestrial air temperature. *International Journal of Climatology*, 15(2), 221229.

CAI	Proposal	Oregon	Precipitation	Climate Interpolation Options
	Climate A	vided Interp	olation (a brie	f review)

Generating daily climate anomalies

$$P_{\text{anomaly}} = \frac{P_{\text{daily}}}{P_{\text{monthlyclimate}}}$$

and temperature:

$$T_{\text{anomaly}} = T_{\text{monthlyclimate}} - T_{\text{daily}} \tag{2}$$

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#### Goal: 1km Global Coverage over 1971-2010

A proposal: shift focus to two products

- 1. 1km Monthly (2000-2011 climatologies): topography + mean monthly station + mean monthly satellite Pros
  - Avoids problems of incomplete daily satellite data (clouds)
  - monthly means much smoother (easier) than daily
  - $\bullet\,$  greatly simplifies processing (12 layers rather than 14,000) Cons

• losing daily information from satellite data

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- losing daily information from satellite data
- 2. Daily climate-aided interpolation for 1970-2011

assume spatial stationarity!



- 1. Oregon Stations
- 2. 70% holdout and 10 sample days identified by Benoit.
- 3. Bayesian 'krige' using spLM in spBayes R package
- 4. compare predictions using climate anomaly and 'raw' station values

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- 5. Simple model with y $\sim$ intercept+lon+lat+elev
- 6. With anomalies, using no covariates fits about the same!

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#### **CAI** Comparison



PRISM climate vs. daily values



Climate Interpolation Options

#### **CAI** Comparison



Note: we should include stations outside ROI to avoid edge effects.

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### CAI Comparison







- 1. Interpolating climate anomalies improves predictive accuracy
- 2. Bayes overall mean RMSE  $2.25^{\circ}C$  (compared with  $\approx 2.39$  for 'best case' other models).

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- 3. Full accounting for predictive uncertainty
- 4. Time consuming (>>2 hours / surface)



Incorporate tmax, tmin, and (perhaps) ppt into a single model?

- 1. Borrow strength across variables (increase n)
- 2. Increases model complexity (and probably run-time)
- 3. Improves fit???

Develop using spBayes package in R http:

//cran.r-project.org/web/packages/spBayes/index.html

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## MODIS Cloud Product (MOD05\_L2)?

Precipitation

Level 2 (swath) data at 1km resolution 2000-2012 Oregon data incudes 3,422 files and  ${\approx}70GB$  data

Layers:

- 1. Cloud Optical Thickness (0 to 100)
- 2. Cloud Optical Thickness Uncertainty (0 to 200%)
- 3. Cloud Effective Radius: particle size (0 to 90  $\mu$ m)
- 4. Cloud Effective Radius Uncertainty(0 to 200%)
- 5. Cloud Water Path (0 to 9000  $g/m^2$ )
- 6. Cloud Water Path Uncertainty (0 to 9000  $g/m^2$ )
- Cloud Phase Optical Properties (0=fill, 1=clear, 2=liquid water cloud, 3=ice cloud, 4=undetermined phase cloud)
- Cloud Multi Layer Flag (0=fill, 1=single layer, 2 through 8=increasing confidence of multilayer clouds)

Partner with TOPS/NEX to get this processed to grid level monthly mean data?



First the climate surfaces:

- 1. MODIS LST (day/max and night/min) monthly means
- 2. MODIS Cloud? TRMM Cloud?
- 3. Use interpolator of choice (spline, krig, GWR)to predict monthly means
- 4. Stationarity: Show past (1970-2000) station data has similar relationship to climate as current (2000-2010)

Then interpolate anomalies:

- 1. GWR? Splines? Kriging? Bayesian Kriging?
- 2. goal changes from explanation (via covariates) to description of pattern



Use existing (continuously updated) gridded anomaly surface to 'add' to our high resolution climate.

Gridded daily GHCN data (2.5° latitude by 3.75°, 1950-present): http://www.metoffice.gov.uk/hadobs/hadghcnd/. Too coarse to be useful??

# Unified Geostatistical Modeling for Data Fusion and Spatial Heteroskedasticity with R Package $ramps^1$

Allows fusion of data obtained at dierent resolutions (areal and point-referenced) and spatial heteroskedasticity.

- 1. Joint modeling of data from multiple sources (point-source, areal, or both).
- 2. Non-spatial random effects as in a linear mixed model.
- 3. Multiple variances for each variation source (measurement error, spatial, and random effects).
- 4. Prediction at measured or unmeasured sites.

Would be useful if we decide to incorporate  $1/4^o$  TRMM data

# Unified Geostatistical Modeling for Data Fusion and Spatial Heteroskedasticity with R Package ramps

$$Y = X\beta + W\gamma + KZ + \epsilon \tag{3}$$

$$\gamma \sim N(0, \Sigma_{\gamma}) \tag{4}$$

$$Z \sim N(0, \Sigma_Z) \tag{5}$$

$$\epsilon \sim N(0, \Sigma_{\epsilon}) \tag{6}$$

where X, W, and K are design matrices for fixed effects  $\beta$   $(p \times 1)$ , non-spatial random effects  $\gamma$   $(q \times 1)$ , and spatial random effects Z  $(S \times 1)$ , respectively.

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